High-pressure phase transitions of ScPO₄ and YPO₄

F. X. Zhang, J. W. Wang, M. Lang, J. M. Zhang, and R. C. Ewin[g*](#page-6-0) *Department of Geological Sciences, The University of Michigan, Ann Arbor, Michigan 48109, USA*

L. A. Boatner

Materials Science and Technology Division and Center for Radiation Detection Materials and Systems,

Oak Ridge National Laboratory, Oak Ridge, Tennessee 37831, USA

(Received 14 July 2009; revised manuscript received 29 October 2009; published 24 November 2009)

 ScPO_4 and YPO_4 with the tetragonal zircon-structure were studied at room temperature and pressures up to \sim 50 GPa. Pressure-induced phase transitions to the sheelite structure occur at 30 GPa for ScPO₄ and 16.3 GPa for YPO4, respectively. In addition to the scheelite-type high-pressure phase, an intermediate phase with the monoclinic monazite-type structure formed during the phase transition process of YPO₄. The high-pressure phases of ScPO₄ and YPO₄ are not quenchable on pressure release. The pressure dependence of the total energy of the different phases was calculated using density-functional method, and the results confirm the experimentally observed phase relations under pressure. Structural parameters and compressibility of each phase were determined by refinement of the x-ray diffraction patterns. The high-pressure phase of ScPO_4 has a very large bulk modulus [376(8) GPa].

DOI: [10.1103/PhysRevB.80.184114](http://dx.doi.org/10.1103/PhysRevB.80.184114)

PACS number(s): 61.50.Ks, 61.66.Fn, 72.80.Ga

I. INTRODUCTION

*AB*O4 ternary oxides are important to a variety of geophysical and geochemical applications, as they are common accessory minerals (i.e., zircon: $ZrSiO₄$; and monazite: $CePO₄$ in a wide variety of rock types in the Earth's upper mantle. $1-4$ $1-4$ The structure-type that forms, i.e., tetragonal zircon vs monoclinic monazite, depends, in part, on the size of the *A*-site cation. For the orthophosphates, both the zircontype and monazite-type structures occur at ambient conditions. The light rare-earth elements (Z less than that of Gd) form the monoclinic monazite-type structure, but the heavier rare-earth elements (including Sc and *Y*) form the tetragonal zircon-type structure. The rare-earth elements at the boundary of these structure types, such as Gd, Tb, and Dy, may occur as both structures at ambient conditions as a function of their thermal history. The structural stability and thermodynamic properties of rare-earth orthophosphates have been systematically investigated[.5](#page-6-3) In both the monazite- and zircon-type phosphates, phosphorous is in fourfold coordination with oxygen forming isolated tetrahedra. The monazite structure is denser because the coordination of the *A*-site cation (including Sc^{3+} and Y^{3+}) is ninefold. These phosphates have excellent high-temperature properties, high radiation damage tolerance and good chemical stabilities. $6,7$ $6,7$ They have been proposed in recent years as host materials for the incorporation and disposal of high-level radioactive nuclear waste. $8-10$ In addition, orthophosphates are of great interest because of their optical properties. $11-13$

Phase transitions of the $ABO₄$ ternary oxides at highpressure have been widely investigated since the 1960's ,¹⁴ and a close relation between cation size and crystal structure was revealed. Fukunaga and Yamaoka¹⁵ have summarized the general trend of the phase transition sequence under pressure as a function of the size of cations. Generally, the radius ratios of r_A/r_O and r_B/r_O between cations and oxygen increase with pressure, resulting in an increase in the cation coordination in the high-pressure phases. A more complete review on various *AB*O4 compounds was completed recently.¹⁶ However, there are rare reports on the pressureinduced phase transitions of rare-earth phosphates. Recently, we have reported the pressure-induced phase transition be-havior in YbPO₄ and LuPO₄,^{[17](#page-6-13)} which is similar with that of zircon at high-pressures. Stavrou *et al.*[18](#page-6-14) have also measured the Raman spectrum of $TbPO₄$ and an irreversible phase transition was reported at pressure of \sim 9.5 GPa, but the structure of the high-pressure phase remains unknown.

 ScPO_4 and YPO_4 are isostructural with the heavy rareearth phosphates, and they both crystallize in the zircon-type structure at ambient conditions. However, their behavior with increasing pressure is different from that of zircon. In this paper, the structural properties of $ScPO₄$ and $YPO₄$ were studied by synchrotron powder x-ray diffraction (XRD), Raman scattering, and theoretical calculations. In addition to the zircon-type to scheelite-type phase transition at highpressure in both systems, an intermediate phase with the monazite-type structure was found during the phase transition process of $YPO₄$. The different structural behaviors of these phosphates are found to be due to the different size of cations.

II. EXPERIMENTAL METHODS

Single crystals of $ScPO₄$ and $YPO₄$ were grown by a high-temperature flux method. Details of the synthesis process were described elsewhere.⁵ In carrying out the XRD measurements, single crystals were crushed into fine powders in a mortar until no obvious separated spots can be found in the diffraction image, and the average grain size was estimated to be less than $2 \mu m$. High-pressure experiments were performed with diamond anvil cell (DAC) techniques and hardened stainless steel gaskets were used for all the high-pressure experiments. Standard methanol/ethanol mixture with or without water was used as the pressure transmitting medium, which maintains good hydrostatic conditions below 10–12 GPa. The *in situ* high-pressure XRD measurements were performed using a synchrotron radiation x-ray source (0.4066 Å) at X17C station of National synchrotron light source (NSLS), Brookhaven National Laboratory, or by using a monochromated x-ray beam (0.4960 Å) at the Cornell high-energy synchrotron source (CHESS). The beam spot size is less than 25×25 μ m² in NSLS and \sim 50 μ m at CHESS. The diffraction rings were recorded with a charge-coupled device (CCD) detector or a Mar345 image plate detector. All of the XRD patterns were integrated from the images using the FIT2D software package.¹⁹ The pressure in all of the experiments was measured by the standard ruby fluorescence method, 20 and the XRD patterns were refined by Rietveld method (Fullprof in the software package Winplotr) 21 21 21 based on the known structural models. In order to obtain the unit cell parameters of the high-pressure phase of ScPO4, an additional XRD experiment was made at 16- ID-D beam of HPCAT, Argonne National Laboratory with monochromated beam of 0.4125 Å and beam size \sim 5 μ m.

Raman measurement was performed at U2 station of NSLS. Small pieces of $ScPO₄$ and $YPO₄$ were loaded in the same pressure chamber. The activate source is 514.3 nm green light from an Ar laser with power of 300 mW. The backscattering light was collected with a liquid nitrogen cooled CCD detector.

Quantum mechanical calculations were performed using the density-functional theory (DFT) framework, and planewave basis sets as implemented in the VASP package. 22 The projector-augmented wave (PAW) method²³ and exchange correlation as parameterized by the Perdew-Wang 91 (PW91) functional^{24[,25](#page-6-21)} were applied in the generalized gradient approximation $(GGA-PAW).^{26}$ $(GGA-PAW).^{26}$ $(GGA-PAW).^{26}$ The total energy and volume calculations for the two systems were completed by optimizing the structures without symmetry constraints. The Monkhorst-Pack scheme for integration in the Brillouin zone is adopted. Test runs were performed to test the energy convergence on *k*-point grid and energy cutoff. A $5 \times 5 \times 5$ *k*-point grid is used for the monoclinic and tetragonal phases, and a $7 \times 7 \times 4$ *k*-point grid is used for high-pressure tetragonal phase. The energy cutoff for the plane-wave basis set is 520.00 eV $(38.22$ Ry) for all the calculations. These parameters are sufficient for energy convergence.

III. RESULTS AND DISCUSSION

A. XRD measurements

The flux grown crystals of $ScPO₄$ and $YPO₄$ are transparent and display good crystal forms. The selected powder XRD patterns of $ScPO₄$ and $YPO₄$ at various pressures are shown in Figs. $1(a)$ $1(a)$ and $1(b)$, respectively. The intensity of the XRD pattern of ScPO_4 is not strong and a small amount of impurities from the residual flux were occasionally observed in the starting material. With the increase in pressure, a new broad peak at two theta around 8.7° appeared clearly in the pattern after 34.2 GPa, which indicates a pressureinduced phase transition. The peak became stronger at higher pressures and more of the sample was transformed to the high-pressure phase(s). However, up to 45 GPa, the zircontype ScPO4 phase is still observed. The broad peak is the

FIG. 1. (Color online) Selected XRD patterns of phosphates at various pressures. (a) SePO_4 measured at NSLS ($\lambda = 0.4066$ Å). Peaks marked with star symbols are due to the residual impurity from the flux during crystal growth. New peaks are clearly observed in the pattern at 34.2 GPa. The peaks marked with up arrows are due to impurity or other minor phase transitions. (b) YPO_4 measured at CHESS $(\lambda = 0.496 \text{ Å})$. The monazite-type high-pressure phase is clearly observable at 16.3 GPa (or even earlier). The pattern after 30 GPa is a combination of monazite- and scheelite-type high-pressure phases.

strongest diffraction maxima of (013) and (112) of the scheelite-type structure. The XRD pattern of $ScPO₄$ at 34.2 GPa was refined as a mixture of the zircon-type and scheelite-type high-pressure phases [Fig. $2(b)$ $2(b)$]. The XRD patterns of the high-pressure phase of $ScPO₄$ were very weak. In order to increase the XRD intensity for ScPO_4 , an additional experiment was performed at HPCAT. The pressure-induced phase transitions measured at HPCAT were similar to the patterns shown in Fig. $1(a)$ $1(a)$. The unit cell parameters of the high-pressure phase of $ScPO₄$ were derived by Rietveld refinement of the patterns measured at HPCAT. In addition to the phase transition to the high-pressure scheelite-type phase, another peak appeared in the highpressure XRD patterns measured at both stations [shown as arrows pointing upwards in Fig. $1(a)$ $1(a)$] at a lower diffraction angle, which could not be indexed as a diffraction peak of either of the two known phases. This peak may originate from some impurities in the sample or be another minor phase transition of $ScPO₄$ at high pressures.

Because of the larger cation size, $YPO₄$ has a much lower critical transition pressure. At 14.6 GPa (or even lower pressures), a new phase is clearly evident from the XRD patterns [Fig. [1](#page-1-0)(b)]. The XRD pattern of the high-pressure phase of YPO4 is very different from the patterns of the high-pressure phases of YbPO₄ and LuPO₄ previously observed.¹⁶ Further analysis revealed that they may be attributed to a lowersymmetry monazite-type phase. The zircon-type $YPO₄$ nearly completely transformed to the monazite-type highpressure phase at pressure above 19.2 GPa because the second strongest peak (112) at two theta angle of 11.2° disappeared completely in the pattern. The XRD pattern of YPO₄ at 16.3 GPa was refined as a mixture of monazite-type and zircon-type phases [Fig. $2(d)$ $2(d)$], and the results indicated a 3.2% volume shrinkage during the transition process. With a

FIG. 2. (Color online) Rietveld refinement results of the observed XRD patterns. (a) ScPO_4 at ambient conditions is the zircontype structure [small amount of Fe (from the steel gasket) and impurities (marked with * symbols)]; (b) XRD pattern of ScPO₄ at 34.2 GPa refined with the mixture of zircon-type and monazite-type phases; (c) YPO₄ at room conditions is a pure zircon-type structure; (d) XRD pattern of YPO₄ at 16.3 GPa refined with the mixture of zircon- and monazite-type structures; and (e) XRD pattern of YPO₄ at 34.9 GPa refined with a mixture of monazite- and scheelite-type structures. The patterns on d and e are measured at CHESS with 0.496 Å x-ray beam and other patterns are measured in NSLS with 0.4066 Å x-ray beam.

further increase in pressure, the profile of the XRD pattern changed at pressures above 30 GPa, suggesting that another phase transition has occurred.

The XRD patterns of $YPO₄$ at pressures higher than 30 GPa are very similar to the patterns of scheelite-type structure, but they contain some extra peaks, such as the triple (or more) peaks at the two theta angle between 5° and 7° . These peaks cannot be attributed to the scheelite-type structure. Analysis of XRD patterns indicated a mixture of two phases for the $YPO₄$ at pressures greater than 30 GPa. The XRD pattern at 34.9 GPa can be refined as a mixture of monazitetype and scheelite-type phases [Fig. $2(e)$ $2(e)$]. Thus, the pressure-induced phase transitions of $ScPO₄$ and $YPO₄$ are different. The zircon-type YPO_4 first transformed to a highpressure monazite-type phase at 14.6 GPa and then to a scheelite-type phase at higher pressures. The phase transitions from zircon-type ScPO_4 and monazite-type YPO_4 to scheelite-type phases are sluggish at room temperature. The sluggish phase-transition behavior at ambient conditions has been previously observed in the studies of zircon $27,28$ $27,28$ and other zircon-type $ABO₄$ compounds,²⁹ because the phase

FIG. 3. (Color online) The XRD patterns of during pressure releasing process for (a) ScPO_4 and (b) YPO_4 . The arrows in Fig. $3(a)$ $3(a)$ show the peaks from pressure medium (water). In the quenched sample, the reflections from the starting zircon-type structure (marked with solid spots) are clearly observed, but the highpressure phases are still detectable in the experimental time scale (especially to $YPO₄$).

transition is displacive, but high temperature greatly enhances the transition. 30

At room temperature, it may be difficult to get a single high-pressure phase for ScPO_4 and YPO_4 . For TbPO₄, we observed a transition sequence similar to that of $YPO₄$, however, up to 60 GPa, the product is still a mixture of the scheelite- and monazite-type phases. The high-pressure phases of $ScPO₄$ and $YPO₄$ are not stable at room conditions. Figure [3](#page-2-1) shows the XRD patterns of $ScPO₄$ and $YPO₄$ during the pressure unloading process. The starting zircon-type phase was recovered in the quenched sample, although some strong diffraction maxima from the high-pressure phases are still observable during the time scale of our experiments.

The unit cell parameters of the different phases in $ScPO₄$ and $YPO₄$ that form as a function of pressure were determined by Rietveld refinement of the XRD patterns. The pressure dependence of the unit cell volume for ScPO_4 and YPO_4 are plotted in Figs. $4(a)$ $4(a)$ and $4(b)$, respectively. The volume decreased about 12.6% at 30 GPa for $ScPO₄$ during the phase transition to the scheelite-type high-pressure phase. For YPO4, the zircon-type phase transformed to monazitetype high-pressure phase at 16.3 GPa with a volume decrease in 3.2% and further transformed to the scheelite-type highpressure phase at 32 GPa with 6.8% volume decrease. Fitting the P-V curves with Birch-Murnagham equation of state yields bulk moduli of 203(7) GPa and 186(5) GPa for the zircon-type $ScPO₄$ and $YPO₄$, respectively. The P-V curve of the corresponding high-pressure phases indicated that they are less compressible. Results obtained by fitting indicated that the monazite-type high-pressure phase of $YPO₄$ has a bulk modulus of 260(29) GPa, when its pressure derivative at zero pressure is fixed at 4. The scheelite-type highpressure phase is a dense form and should have larger bulk modulus. The two data points for scheelite-type $YPO₄$ do not give a reliable value for the bulk modulus. Fitting the volume data for scheelite-type $ScPO₄$ yielded a very large bulk modulus of 376(8) GPa $(B'_0=4)$. This is even greater than

FIG. 4. (Color online) Pressure dependence of the unit cell volume of (a) $ScPO_4$ and (b) YPO_4 . There is a volume collapse of \sim 12.6% at 30–34 GPa for ScPO₄ during the phase transition from zircon-type to scheelite-type high-pressure phase. The volume collapse for $YPO₄$ is 3.2% at 16.3 GPa for the phase transition from zircon-type to monazite-type and 8.2% at 32.3 GPa from the monazite-type to scheelite-type.

that of reidite $(ZrSiO₄)$ —the high-pressure form of zircon (301 GPa) ,^{[31](#page-6-27)} and much larger than the values of scheelitetype $YbPO₄$ and $LuPO₄$.^{[17](#page-6-13)} The large bulk modulus suggests that the high-pressure form of $ScPO₄$ may be a superhard material. Unfortunately, the scheelite-type high-pressure phase of ScPO_4 could not be quenched to room conditions.

The bulk modulus of silicates and some binary oxides was found to be directly correlated to the compressibility of the *A*-cation coordination polyhedra[,32](#page-6-28) and it obeys a linear relationship to the compressibility of AO_8 polyhedron in the following empirical formula³³ to most zircon-related $ABO₄$ structures,

$$
B_0 = 610(110)\frac{z_A}{d_{A-O}^3},\tag{1}
$$

where B_0 is the bulk modulus in GPa, z_A is the formal charge of *A* cations ($1 \le z_A \le 4$), and d_{A-O} is the average *A*-O distance (\AA) in the $A\text{O}_8$ polyhedron. To most $AB\text{O}_4$ compounds, *B*O4 tetrahedra are interconnected through O-*A*-O bonds, and the B -O bonds in $BO₄$ tetrahedra are shorter and stronger than the A -O bond in A O₈ polyhedra. As a result, the compressibility of the whole lattice is mainly related to the compressibility of AO_8 polyhedra. The linear relationship between bulk modulus and the charge density of *A* cation is consistent with the fact that AO_8 polyhedra exhibiting a larger electronic cloud inside the polyhedral than those AO_8 with a lower *A*-cation charge density. The predicted bulk modulus with the above formula fit the observations to most silicates and transition metal oxides.^{16[,32](#page-6-28)} However, the predicted bulk modulus of phosphates are obvious lower than the experimental results, especially to the scheelite-type structures.¹⁷ At normal conditions, the bonding distances of Sc-O and Y-O in zircon-type ScPO_4 and YPO_4 are 2.2016 and 2.3365 Å, respectively, and the predicted bulk modulus are 171 and 144 GPa, which are less than the observed values (203 GPa for $ScPO₄$ and 186 GPa for $YPO₄$). In phosphates, cation P^{5+} is four-coordinated and there is an extra electron transfer from the $P^{5+}O_4$ tetrahedron to the neighbor A_0 ₈ polyhedra, and as a result, the effected electron density in $AO₈$ polyhedra increased. If the effective charge of *A* cation is taken as 3.6, the calculated bulk modulus is 205 GPa for the zircon-type ScPO_4 and 169 GPa for YPO₄, which fit the experimental results quite well. The scheelite-type structure is a dense form of the zircon-type structure, and the AO_8 polyhedra are compressed. For example, the bond distance of Zr-O in reidite (scheelite-type zircon, $ZrSiO₄$) is 2.22 Å, which is smaller than that in zircon $(2.268$ the bond distance of Zr-O), while the Si-O bond distance in reidite is 1.68 Å and larger than that in zircon (1.525 Å) . The bulk modulus is related to both AO_8 polyhedra and BO_4 tetrahedra when their sizes are comparable, and the linear relationship stated above will not be applicable. There are obvious deviation between the observed and predicted bulk modulus for zircon $(209 \text{ GPa} \text{ calculated from formula } (1) \text{ and more than } 300$ $(209 \text{ GPa} \text{ calculated from formula } (1) \text{ and more than } 300$ $(209 \text{ GPa} \text{ calculated from formula } (1) \text{ and more than } 300$ GPa for observed).^{[31](#page-6-27)} In scheelite-type phosphates, Sc has the smallest size (the bond distance is even smaller than Zr-O in reidite), so the compressibility should be more relied on both ScO_8 polyhedra and PO_4 tetrahedra, and the very high bulk modulus is thus understandable.

B. Raman measurements

Raman measurement of zircon and the assignment of the individual modes have been well done defined before. $34,35$ $34,35$ Recently, a similar result on zircon-type $TbPO₄$ was reported based on polarization measurements of single crystals.¹⁸ Group theory analysis predicts that there are 12 Raman active modes in the Brillouin zone center:

$$
\Gamma = 2A_{1g} + 4B_{1g} + B_{2g} + 5E_g.
$$
 (2)

These modes can be classified as seven internal modes $(2A_{1g} + 2B_{1g} + B_{2g} + 2E_g)$, four external translational modes $(2B_{1g} + 2E_g)$, and one external rotational mode (E_g) . The internal modes are due to the normal vibrations of oxygen atoms within the $PO₄$ tetrahedra. The translational modes are generated from the translations of $PO₄$ tetrahedra and the larger cations (Sc³⁺, Y³⁺, or Ln³⁺). The rotational mode E_{g} represents the rotation of whole $PO₄$ tetrahedral units.

Crystals of $ScPO₄$ and $YPO₄$ have very good Raman signals at ambient conditions, and nearly all the 12 Raman active modes can be detected. However, with the increase in pressure, the signal/noise ratio decreases dramatically, and no good signal can be detected at pressure higher than 15 GPa for the small pieces of both crystals $(< 10 \mu m)$. The measured Raman spectra of $ScPO₄$ and $YPO₄$ at various pressures are shown in Fig. [5.](#page-4-0) Above 10 GPa, the Raman signal from both samples became very weak. For ScPO_4 , the Raman measurement could not be done at the pressure range of the transition. The Raman modes of $YPO₄$ above 13.4 GPa became very weak, and even the strongest A_{1g} mode $($ ~1000 cm⁻¹ at ambient conditions) could not be detected. However, a peak with a frequency a little lower than that of the A_{1g} mode appeared [marked with arrow in Fig. $5(b)$ $5(b)$], which indicated a pressure-induced phase transition in YPO₄. The smaller value of the mode from the high-pressure phase of YPO_4 is consistent with the observations of TbPO₄.^{[18](#page-6-14)} The measured phase transition of YPO₄ by Raman occurred at slightly lower pressures (14.6 GPa) than observed by XRD measurements. This is reasonable because Raman scattering

FIG. 5. (Color online) Raman spectrum of phosphates measured at different pressures (a) SePO_4 and (b) YPO_4 . A phase transition can be observed in $YPO₄$ at 13.4 GPa, and the mode marked with arrow is contributed by high-pressure phase.

measures the short-range order, thus, a local structural change can be detected at high pressures. The pressure dependence of the measured wave number of the vibrational modes for ScPO_4 and YPO_4 are shown in Fig. [6.](#page-4-1) The assignment of individual modes is based on the polarization mea-surements in TbPO₄.^{[18](#page-6-14)} The assignments, wave numbers of individual modes at ambient conditions, and their pressure dependence are listed in Table [I.](#page-4-2) The Gruneisen parameters of the individual modes of the zircon-type phase were calculated with the formula

$$
\kappa = \frac{B_0}{\omega_i} \frac{\partial \omega_i}{\partial P},\tag{3}
$$

where ω_i is the Raman wave number of the mode, and B_0 is the bulk modulus from XRD measurements. Importantly, there are two sharp modes—one external translational mode,

FIG. 6. (Color online) Pressure dependence of the wave number of different modes in (a) SePO_4 and (b) YPO_4 . The open symbols represent the soft modes.

 $E_{\rm g}$, and one internal mode, B_{2g} , which are softened at high pressures, and their wave numbers show a slight decreases with the increase in pressure. The rotational mode in the phosphate units of ScPO_4 and YPO_4 has the largest shift at high-pressure. This is very different from the behavior of polyhedra at elevated pressures in other solids, such as the boron-icosahedra in boron-rich solids.^{36,[37](#page-6-33)} In boron-rich solids, all of the B_{12} icosahedra lie along lines with the same orientation, and the compression of the structure has little effect on the rotation of the polyhedral around these lines; thus, the rotational mode is nearly pressure independent. In these phosphates, however, the $PO₄$ tetrahedra are separated from one another by larger cations $(Sc³⁺ or Y³⁺)$, and they do not lie along the same axis or have the same orientation, and thus be easily affected by external pressure.

C. Theoretical calculations

Previous experimental results revealed that the critical transition pressure and structural stability of $ABO₄$ com-

TABLE I. Raman active modes of ScPO_4 and YPO_4 at room conditions, assignments and their pressurerelated Grunium parameters (κ) .

		ScPO ₄		YPO ₄	
Character	Assignments	ω_0 (cm^{-1})	К	ω_0 (cm^{-1})	K
External translational	$E_{\rm g}$	186.7	-0.64	157.1	-0.79
	B_{1g}	233.7	0.74	183.6	1.18
	$E_{\rm g}$ (or $B_{1{\rm g}}$)	244.4	1.22	210.0	0.65
Rotational	$E_{\rm g}$	335.4	2.86	297.9	3.16
Internal	B_{2g}	325.3	-0.25	330.1	-0.31
	A_{1g}	474.3	0.41	482.5	0.43
	$E_{\rm g}$	594.4	0.44	579.3	0.40
	B_{1g}			656.4	0.58
	A_{1g}	1027.2	0.80	999.6	0.97
	$E_{\rm g}$	1044.5	0.75	1025.4	0.90
	B_{1g}	1083.0	0.78	1057.0	0.94

FIG. 7. (Color online) Calculated lattice energy of the different phases in (a) ScPO_4 and (b) YPO_4 .

pounds are strongly relied on the size ratio of $BO₄$ tetrahedra and *A* cation.¹⁶ Zircon-type structure is more stable to those compounds with larger values of size ratio. To rare-earth phosphates, only *A* cations with smaller size (including Sc and Y) can form zircon-type structure, and all the larger cations from La and Gd form the monazite-type structure at ambient conditions. Gadolinium phosphate is on the boundary and the critical value of the size ratio of BO_4/A is 2.65. Zircon-type phase is stable at ambient condition to those phosphates with PO_4/A size ratio larger than 2.65. Due to the small size of Sc, ScPO₄ has the largest size ratio (\sim 3.3) and is thus the most stable zircon-type phosphate. In zircon-, scheelite- and monazite-type structures, rare-earth (RE) are loosely bonded with 8 or 9 neighbor anions. Under highpressure, the RE-O bonds are easier to be compressed, compared to the strong bonded $PO₄$ tetrahedra. The competition and stability of the above three phases at high pressures can only be determined by quantum mechanical calculations. Our in situ XRD experiments under high-pressure for all rare-earth phosphates indicated that only Yb and Lu (also Sc) can form pure scheelite-type structures directly from the zircon-type structures at high pressures.¹⁷ Cation Y^{3+} has similar size with Dy^{3+} and Ho^{3+} , and experiments found that zircon-type phosphates with rare-earth elements from Gd to Tm has a pressure-induced phase transition from zircon-type

to monazite-type phase first and then tend to transform to the scheelite-type structure. The formation of intermediate monazite-type structure in $YPO₄$ at high-pressures is thus expected.

In order to further understand the different phase transition behaviors of $ScPO₄$ and $YPO₄$ at high-pressure, a quantum mechanical calculation was made based on densityfunctional theory with plane-wave basis sets using VASP. The structure of the zircon-type and the scheelite-type highpressure phases for both samples are based on the known structural model of zircon and the XRD measurements. The initial structure of the monazite-type high-pressure phase is built on the monoclinic phase of LaPO₄. The calculated total energy as a function of volume of different phases of ScPO₄ and $YPO₄$ is shown in Fig. [7.](#page-5-0) The calculated bulk moduli are listed in Table [II](#page-5-1) together with the measured values. The calculated values are in general agreement with our experimental observations. The calculated phase transition sequence also indicates that a monazite-type high-pressure phase of $YPO₄$ forms first, and the scheelite-type highpressure phase becomes more stable at higher pressures. The calculated transition pressures from zircon-type to monazitetype YPO_4 is $P_1 = 15.8$ GPa, which is in good agreement with the observed value of 14.6 GPa. The calculated transition pressures to the scheelite-type, high-pressure phase in both samples $(21.7 \text{ GPa} \text{ for } \text{ScPO}_4$; 20.8 GPa for YPO_4) are a little less than the measured transition pressures (both greater than 30 GPa). The calculated monazite-type phase of ScPO4 has a slightly smaller volume as compared with the zircon-type structure. The phase transition from zircon-type to monazite-type is difficult to induce by pressure at low temperature. The monazite-type structure of $ScPO₄$ may be a high-temperature phase. There are no reports of phases other than zircon-type structure in ScPO_4 and YPO_4 up to now.

IV. SUMMARY

Pressure-induced phase transitions of zircon-type ScPO₄ and YPO₄ were investigated by powder XRD, Raman scattering, and quantum mechanical calculations. The zircontype phosphates transformed to scheelite-type high-pressure phases at pressures above 30 GPa in $ScPO₄$ and $YPO₄$. There

TABLE II. Measured and calculated bulk modulus of various structures of SePO_4 and YPO_4 The measured bulk modulus is derived from the fitting of P-V curve with Birch-Murnaghan equation of state when the derivative B_0' is fixed at 4. The theoretical values are derived from the calculated energy-volume curves and pressure calibration was applied to the calculated P-V data before the bulk modulus was calculated.

		SePO _A		YPO _A		
Structure	Measured	Calculated	Measured	Calculated		
Zircon type	203 (7) GPa $(0-35 \text{ GPa})$	183.0 (2.3) GPa B'_0 = 4.2(0.2)	186 (5) GPa $(0-21 \text{ GPa})$	165.5 (1.4) GPa B'_0 = 4.3(0.1)		
Monazite type		188.5 (0.1) GPa $B'_0 = 3.3(0.1)$	260 (29) GPa $(16-34 \text{ GPa})$	190 (13) GPa B'_0 = 2.4(3.7)		
Scheelite type	376 (8) GPa $(31 - 50 \text{ GPa})$	334 (16) GPa $B'_0 = 3.4(0.2)$		213.7 (1.1) GPa $B'_0 = 3.6(0.1)$		

is a monazite-type high-pressure phase formed before the transition from zircon-type to scheelite-type structure in $YPO₄$ at 14.6 GPa. All the sluggish phase transitions in $ScPO₄$ and $YPO₄$ are displacive and reversible. The observed phase transition sequences and transition pressures for the two systems are due to the different cationic size, and consistent with the results of the quantum mechanical calculations. The very large bulk modulus of the scheelite-type high-pressure phase of $ScPO₄$ suggests that it may be a superhard material.

ACKNOWLEDGMENTS

This work was supported by the Office of Basic Energy Sciences of the U.S. Department of Energy, through Grant No. DE-FG02-97ER45656. Research at ORNL is sponsored by the Division of Materials Science and Engineering, Office of Basic Energy Sciences, US DOE, under Contract No. DE-AC05-00OR22725 with Oak Ridge National Laboratory, managed and operated by UT-Battelle, LLC. The use of the beam line at X17C and U2 station of NSLS is supported by NSF COMPRES under Contract No. EAR01-35554 and by US-DOE Contract No. DE-AC02-10886. The use of the synchrotron in the B2 station of CHESS at Cornell University is financially supported by the National Science Foundation and the National Institutes of Health/National Institute of General Medical Sciences under NSF Grant No. DMR-0225180. The synchrotron experiments performed at HPCAT (Sector 16), Advanced Photon Source (APS), Argonne National Laboratory is supported by DOE-BES, DOE-NNSA, NSF, and the W.M. Keck Foundation. APS is supported by DOE-BES under Contract No. DE-AC02-06CH11357.

*rodewing@umich.edu

- 1R. J. Finch and J. M. Hanchar, Rev. Mineral. Geochem. **53**, 1 $(2003).$
- 2A. F. Reid and A. E. Ringwood, Earth Planet. Sci. Lett. **6**, 205 $(1969).$
- 3F. X. Zhang, V. Pointeau, L. C. Shuller, D. M. Reaman, M. Lang, Z. Liu, J. Hu, W. R. Panero, U. Becker, C. Poinssot, and R. C. Ewing, Am. Mineral. **94**, 916 (2009).
- 4M. Lang, F. X. Zhang, J. Lian, C. Trautmann, R. Neumann, and R. C. Ewing, Earth Planet. Sci. Lett. **269**, 291 (2008).
- 5S. V. Ushakov, K. B. Helean, A. Navrotsky, and L. A. Boatner, J. Mater. Res. **16**, 2623 (2001).
- ⁶ J. B. Davis, D. B. Marshall, R. M. Housley, and P. E. D. Morgan, J. Am. Ceram. Soc. **81**, 2169 (1998).
- ⁷ J. B. Davis, D. B. Marshall, and P. E. D. Morgan, J. Eur. Ceram. Soc. **20**, 583 (2000).
- 8R. C. Ewing and L. Wang, Rev. Mineral. Geochem. **48**, 673 $(2002).$
- 9L. A. Boatner, G. W. Beall, M. M. Abraham, C. B. Finch, P. G. Huray, and M. Rappaz, Bull. Am. Phys. Soc. 25, 186 (1980).
- ¹⁰ J. G. Pepin, E. R. Vance, and G. J. McCarthy, Mater. Res. Bull. 16, 627 (1981).
- ¹¹A. J. Wojtowicz, D. Wiskiewski, A. Lempicki, and L. A. Boatner, Radiat. Eff. Defects Solids 135, 305 (1995).
- 12A. Lempicki, E. Berman, A. J. Wojtowicz, M. Balcerzyk, and L. A. Boatner, IEEE Trans. Nucl. Sci. 40, 384 (1993).
- 13W. W. Moses, M. J. Weber, S. E. Derenzo, D. Perry, P. Berdahl, and L. A. Boatner, IEEE Trans. Nucl. Sci. 45, 462 (1998).
- 14F. Dachille and L. S. D. Glasser, Acta Crystallogr. **12**, 820 $(1959).$
- 15O. Fukunaga and S. Yamaoka, Phys. Chem. Miner. **5**, 167 $(1979).$
- 16D. Errandonea and F. J. Manjon, Prog. Mater. Sci. **53**, 711 $(2008).$
- 17F. X. Zhang, M. Lang, R. C. Ewing, J. Lian, Z. W. Wang, J. Hu,

and L. A. Boatner, J. Solid State Chem. **181**, 2633 (2008).

- 18A. T. E. Stavrou, Y. C. Boulmetis, A. G. Kontos, Y. S. Raptis, and C. Raptis, J. Phys.: Condens. Matter **20**, 395222 (2008).
- ¹⁹A. P. Hammersley, FIT2D, ESRF, Grenoble, France, 1998.
- 20H.-K. Mao, J. Xu, and P. M. Bell, J. Geophys. Res. **91**, 4673 $(1986).$
- ²¹ J. Rodriguez-Carvajal, Fullprof 2k, France, 2001.
- ²² J. Hafner, Comput. Phys. Commun. **177**, 6 (2007).
- ²³ P. E. Blöchl, Phys. Rev. B **50**, 17953 (1994).
- ²⁴ J. P. Perdew, J. A. Chevary, S. H. Vosko, K. A. Jackson, M. R. Pederson, D. J. Singh, and C. Fiolhais, Phys. Rev. B **46**, 6671 $(1992).$
- ²⁵ J. P. Perdew and Y. Wang, Phys. Rev. B **45**, 13244 (1992).
- ²⁶ J. P. Perdew and Y. Wang, Phys. Rev. B **33**, 8800 (1986).
- 27 W. van Westrenen, M. R. Frank, J. M. Hanchar, Y. Fei, R. J. Finch, and C.-S. Zha, Am. Mineral. **89**, 197 (2004).
- ²⁸ E. Knittle and Q. Williams, Am. Mineral. **78**, 245 (1993).
- 29A. Grzechnik, K. Syassen, I. Loa, M. Hanfland, and J. Y. Gesland, Phys. Rev. B 65, 104102 (2002).
- 30S. Ono, K. Funakoshi, Y. Nakajima, Y. Tange, and T. Katsura, Contrib. Mineral. Petrol. **147**, 505 (2004).
- 31H. P. Scott, Q. Williams, and E. Knittle, Phys. Rev. Lett. **88**, 015506 (2001).
- ³² R. M. Hazen and C. T. Prewitt, Am. Mineral. **62**, 309 (1997).
- 33D. Errandonea, J. Pellicer-Porres, F. J. Manjón, A. Segura, Ch. Ferrer-Roca, R. S. Kumar, O. Tschauner, P. Rodríguez-Hernández, J. López-Solano, S. Radescu, A. Mujica, A. Muñoz, and G. Aquilanti, Phys. Rev. B 72, 174106 (2005).
- 34P. Dawson, M. M. Hargreave, and G. R. Wilkinson, J. Phys. C **4**, 240 (1971).
- 35A. Gucsik, M. Zhang, C. Koeberl, E. K. H. Salje, S. A. T. Redfern, and J. M. Pruneda, Mineral. Mag. 68, 801 (2004).
- 36D. R. Tallant, T. L. Aselage, A. N. Campbell, and D. Emin, Phys. Rev. B 40, 5649 (1989).
- ³⁷ F. X. Zhang, Chem. Phys. Lett. **379**, 47 (2003).